Seismic imaging of variable water layer sound structure in Rockall Trough, NE Atlantic:

Implications for planning seismic surveys in deep water



Ciaran Sutton, Stephen Jones, Robert Hardy and David Hardy

Seismic and Basin Analysis Group, Department of Geology, School of Natural Science, Trinity College Dublin, Ireland Contact: csutton@tcd.ie

1. Introduction

- Variation in sound speed within ocean water can cause serious problems in subsurface seismic reflection imaging. Water layer sound speed variations arise from natural variations in temperature and salinity associated with oceanic currents.
- In 3D seismic surveying, water layer sound speed variations cause TWT offsets and amplitude variations between sail lines. Vertical sound speed gradients cause ray bending and lead to migration problems.
- 4D seismic surveying compounds these problems.
- Even in a 2D survey, water layer sound speed variation can adversely affect line ties and merging of re-shot sections of a single line.
- Important questions for the hydrocarbon industry include: 1) How might water layer sound speed vary during acquisition of a single sail line?
- 2) How might sound speed vary between acquisition of adjacent sail lines?

3) How might sound speed vary between repeated 3D survey acquisitions in a 4D imaging program?

4) What acquisition and processing strategies are required to cope with the expected variability?

The deep water basins offshore west of the UK and Ireland provide a natural laboratory to study the effect of water layer variability in seismic reflection images, and to compare the signatures of water layer variability in legacy oceanographic and seismic reflection datasets (Fig. 1).



ure 1. Location map of the researc ea West of Ireland. Heavy black lines present the seismic lines shown.

Acknowledgements

All seismic processing was performed using the Seismic Unix system

We thank the Irish petroleum Infrastrucutre Project for supporting our seismic processing facility and the Irish Marine Institute for access to oceanographic datasets. The MOST project currently involves Trinity College Dublin, University of Durham, Memorial University of Newfoundland, GX Technology, Spectrum, Statoil Ireland Exxon-Mobil & BG Group

We thank the Irish Petroleum Affairs Division for highlighting this project to participating oil companies

from Rockall Trough, NE Atlantic, illustrate both pronounced vertical layering and significant space-time variation.

igure 2. Oceanic structure of Rockall pugh. At depths shallower than 1.5km mperature and salinity vary strongly nd generate significant sound speed

deeper water (Fig. 3b).

peed, c)+d) Seabed vertical offsets, e)+f abed lateral shift

- are more likely in deeper water.
- Ray bending can cause many water with vertical sound speed is done using a constant water of rays can significantly impair subsurface images. In Rockall occur in water depths shallower than 1 km, and this problem is equally severe in summer and winter. The mis-positioning problem is less severe in deeper water because of opposing sound speed gradients at different levels in the water column.
- vary at a given location.
- data.
- climate change.



Sound Speed (m/s)

‡ISROCK Survey

Hay-Aug, N=37 N=37 3.0

2. Oceanographic Data

4. North Atlantic

Deep Water

2.5 Rockall Trough

3.0 🗍 May-Aug, N=112

All year, N=181

5 10 15

Salinity (psu

Legacy oceanographic data

Mean water layer sound speed in Rockall Trough mostly varies between 1490 and 1500 m/s (Fig. 3a). Variability is greater in winter and less, though still significant, in summer. Strong variability in mean sound speed occurs above 2 km and variability decreases somewhat in

Figure 3. a)+b) Mean water layer sound



1490 1500 1510

Seabed TWT vertical offsets of over 15 ms are expected if the water layer is assigned a constant sound speed. Misties shown in Fig. 3c & d result when the water layer is assigned the modal mean sound speed of 1496.5 m/s.Poor choices for mean water layer sound speed will lead to larger mis-ties. Large mis-ties

hundred metres of lateral shift in gradients (Fig. 3e & f). If migration layer sound speed, mis-positioning Trough, the largest lateral shifts



gure 4. Four CTD profiles (near Westline Fig. 8) highlighting that sound speed eral variability within layer 2+3 (Fig. 2) occurs at both a monthly and yearly ale, equivalent to the variation in reflections seen in seismic profiles (Section 3).

5. Conclusion

Risk assessment for water layer variation is essential in 3D and 4D seismic acquisition and processing. Risk also affects line ties and re-shot segment ties in 2D surveys.

Use legacy oceanographic data for initial assessment of water layer risk. Use legacy seismic data to refine the risk assessment by showing how rapidly and by how much the water layer sound speed is likely to

In areas of vigorous oceanic mixing, such as offshore west of the UK and Ireland, expect strong water layer variability along individual sail lines. Pick average water layer speeds at a fine horizontal scale (< 1km).

In areas of strong vertical sound speed gradients, it is necessary to account for ray bending when migrating

Make the water layer of seismic data avialbale for academic research into oceanic mixing. Improved understanding of mixing will feed through into improved models of global oceanic circulation and global

3. Seismic data

- Ocean data indicate strong oceanic variability on a basin scale. What is the variability on the space-time scale of a seismic survey?
- In Rockall Trough, the oceanic variability is clearly visible as a strongly reflective layer at depths between 500 m and 1.5 km. Related to strongly developed temperature and salinity fine structure observed over this depth range in the oceanographic data.
- Strong lateral variations of reflectivity are visible. Some features can be interpreted as eddies originating in Mediterranean outflow water and carried into Rockall by the along-slope current. The corresponding thermo-haline variations are expected to be associated with strong sound speed variations.
- Some reflective packages have steep and sharp edges. Such edges will generate abrupt changes in average sound speed along individual seismic sail lines. Figures 6-8 show these features.

5 Line 64, 56, 54, 52 from the ISROCK96 survey NE Rockall, g reflections in the water column that correspond to ges in sound speed





Figure 8 Westline NW-SE Rockall

Westline

- Re-shot sections (owing to gun failure) show that there can be significant variability in water layer reflectivity at a single location over a matter of hours.
- These reflections are the result of variations in temperature and salinty. Over time these water properties change and therefore different reflections are observed over the 3-5 hour reshoot gap. These reflections can be quantified in terms of average seismic energy.

gure 9. a) Line9 b) Line58 c) Line60 d) Line74 and responding reshoot. Plotted on top is the average seismic energy for the water column



witer witer strok-58-50 Time

East

