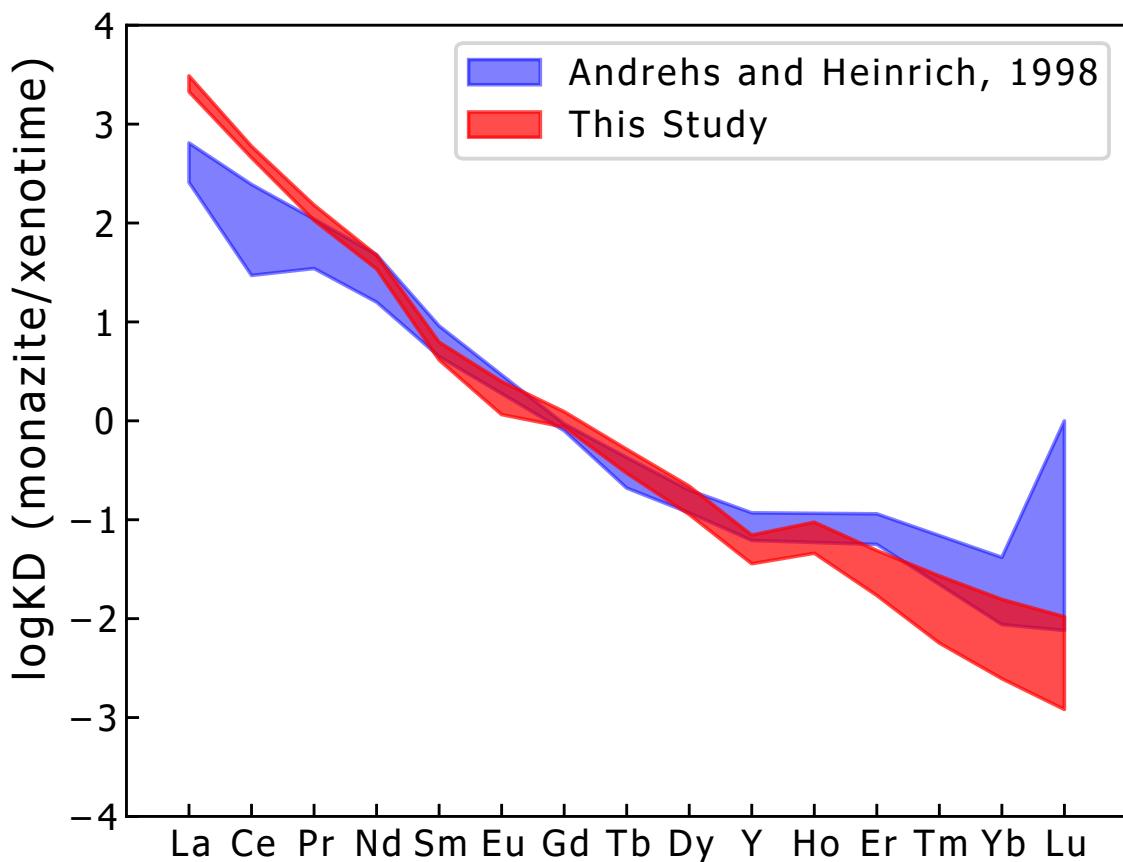
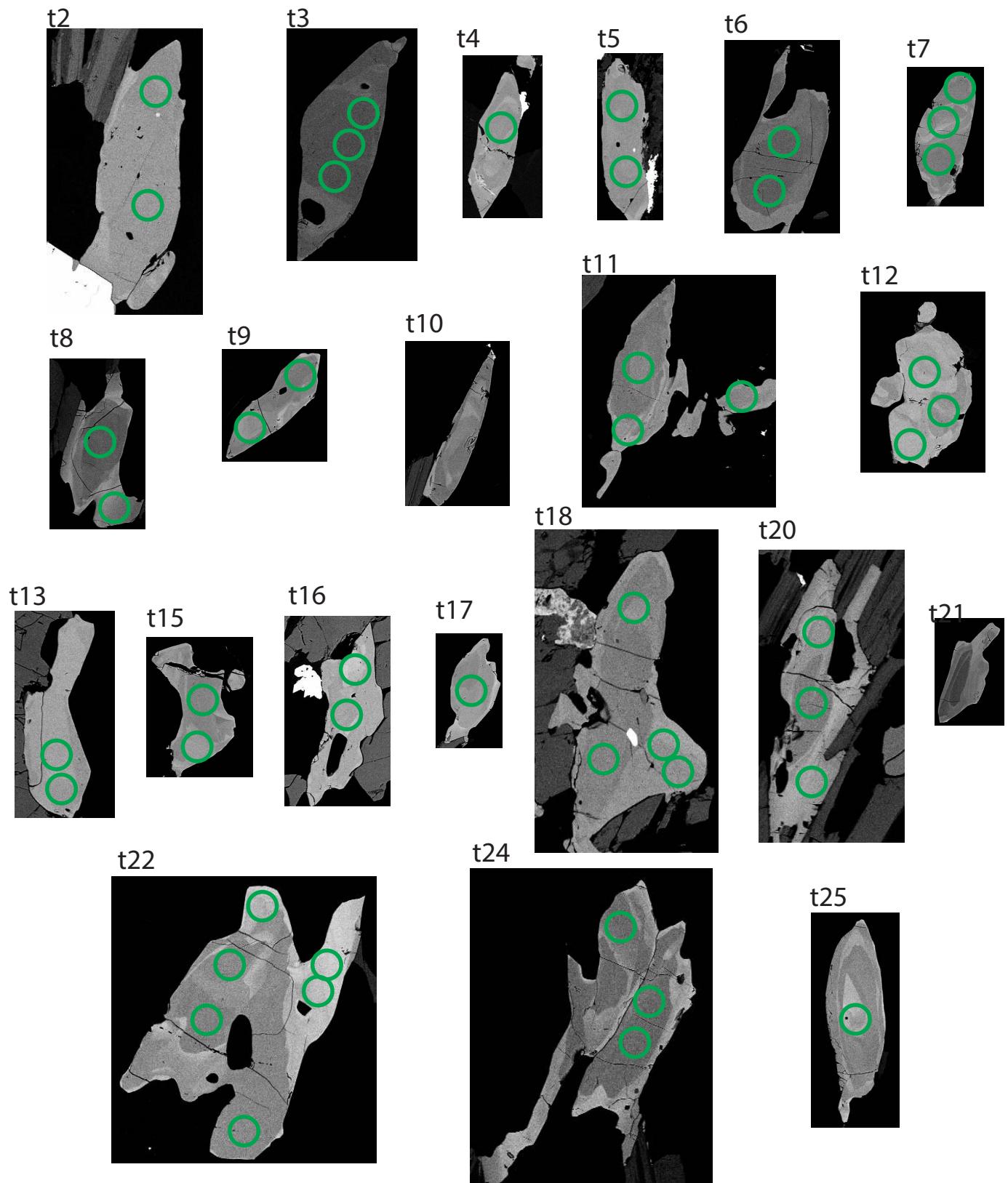


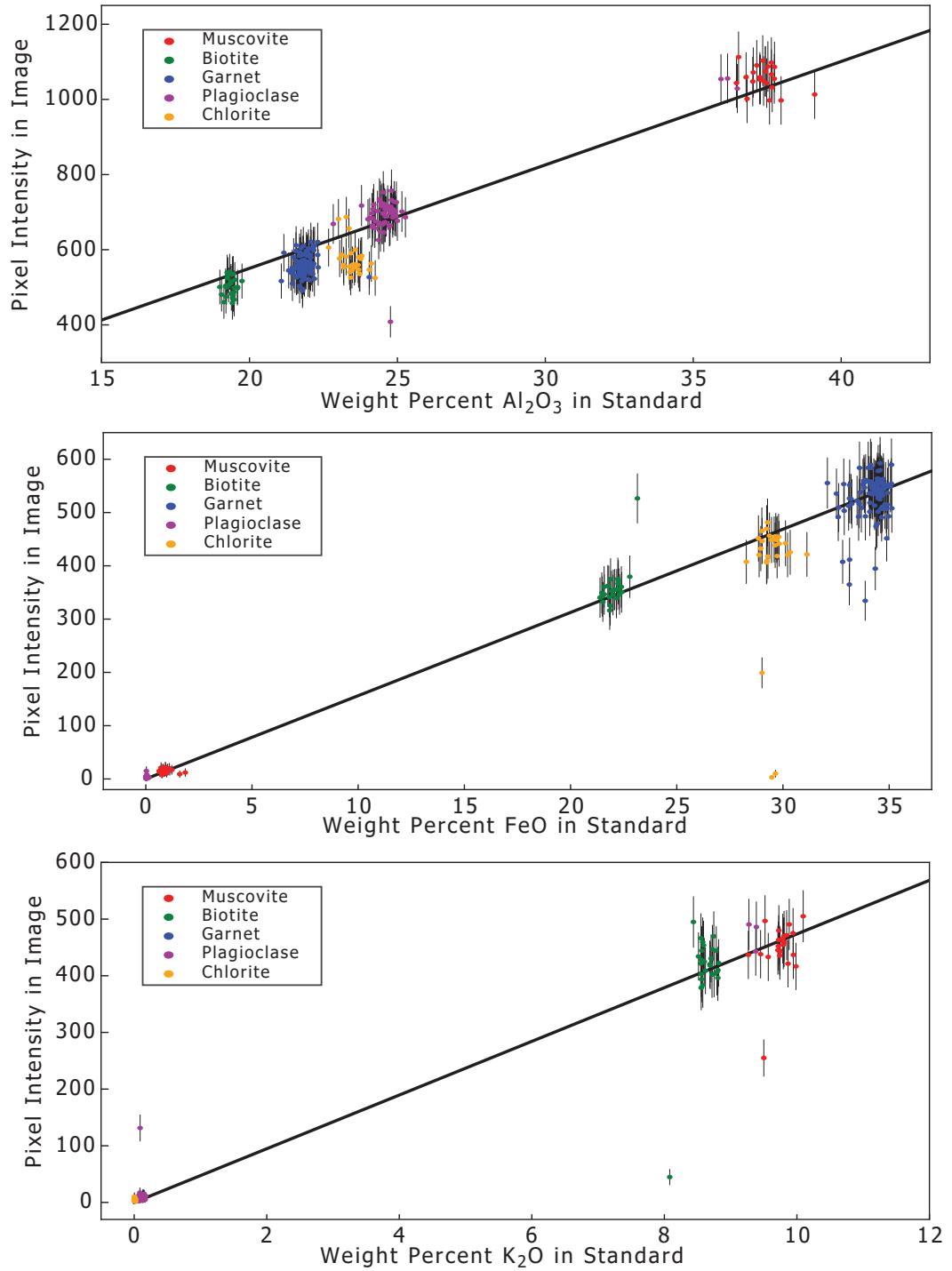
Supplementary Figure 1: Terra-Waserburg diagram showing data from monazite (blue) and xenotime (red) older than 500 Ma. These data may relate to older metamorphic events such as the the Knoydartian orogeny. 930-940 Ma monazite (included in garnet) may be detrital in origin.



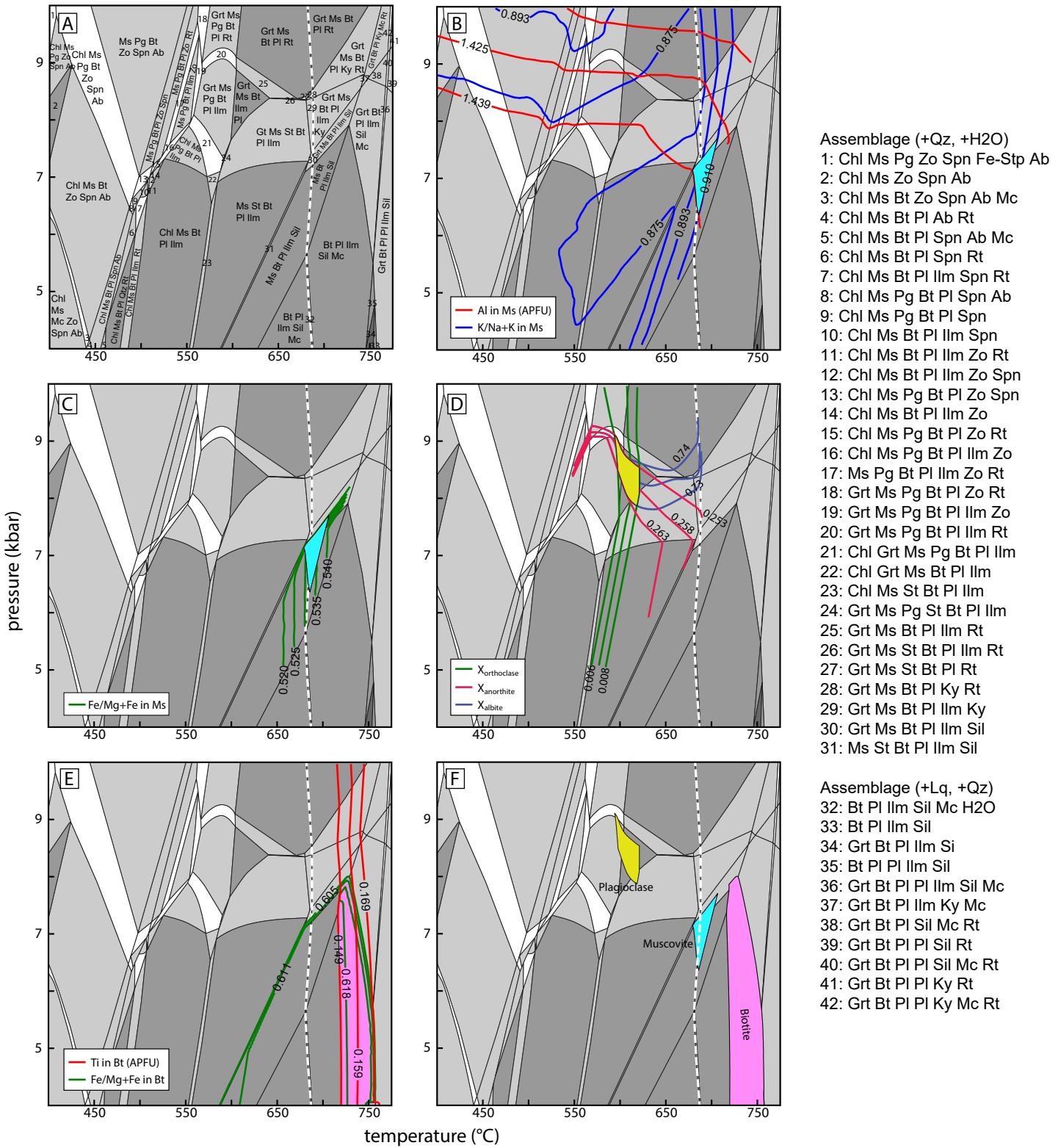
Supplementary Figure 2: A comparison of REE partitioning in monazite and xenotime from this study and that of Andrehs and Heinrich (1998). The range of our data mostly overlaps with the experimental data of Andrehs and Heinrich (1998). The general match can be regarded as an indicator of equilibrium between monazite and xenotime, so greater confidence can be placed in the constraints from monazite-xenotime thermometry.



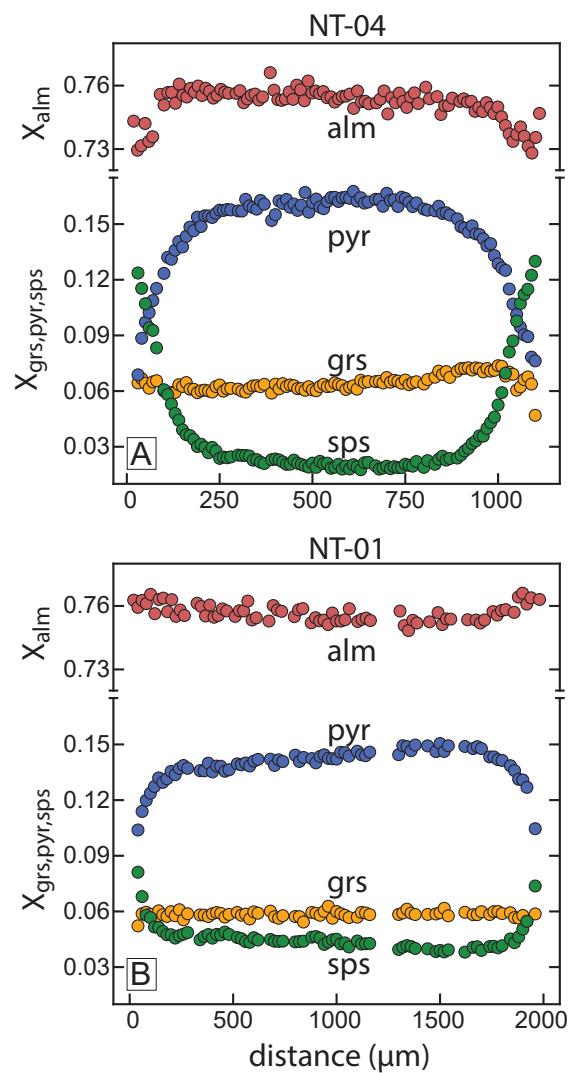
Supplementary Figure 3: Backscatter electron images of titanite with locations of laser ablation spots. These can be associated with isotopic and trace element data using x-y coordinates of titanite spots listed in supplementary table 2.



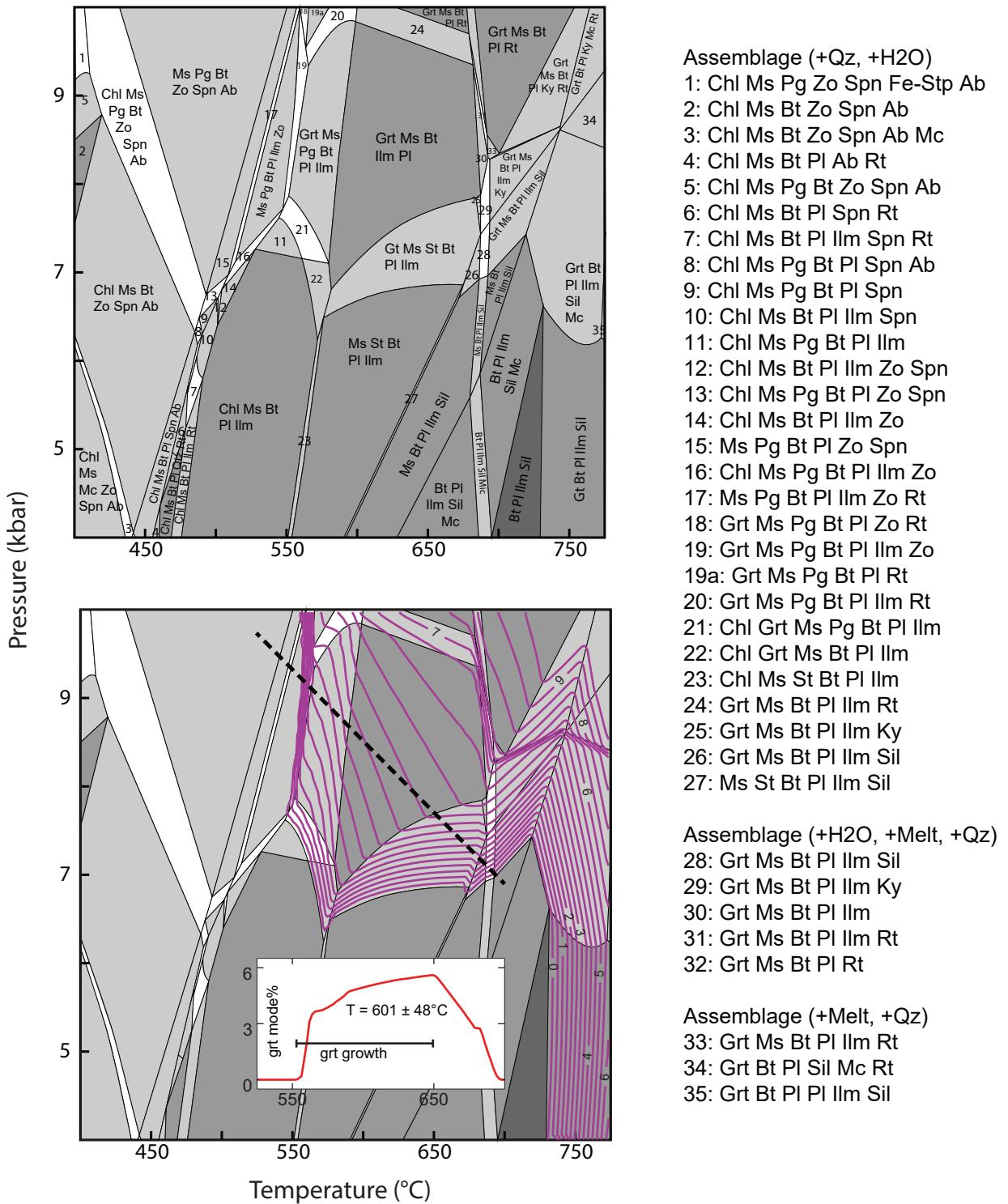
Supplementary Figure 4: Calibration of x-ray images using XMapTools (Lanari et al., 2014). The relationship between EPMA-measured composition and pixel intensity (counts per second) in an x-ray image of the same area is used to calculate the composition of the whole image area. These calibrations were used to calculate the compositions of coarse-grained muscovite in sample NT-04 (NT-14-01), which was not measured by EPMA, but was captured in x-ray images. We used EPMA data from all mineral phases to construct the calibration. This approach appears to give much more reasonable muscovite compositions than are obtained from a calibration using only muscovite EPMA analyses.



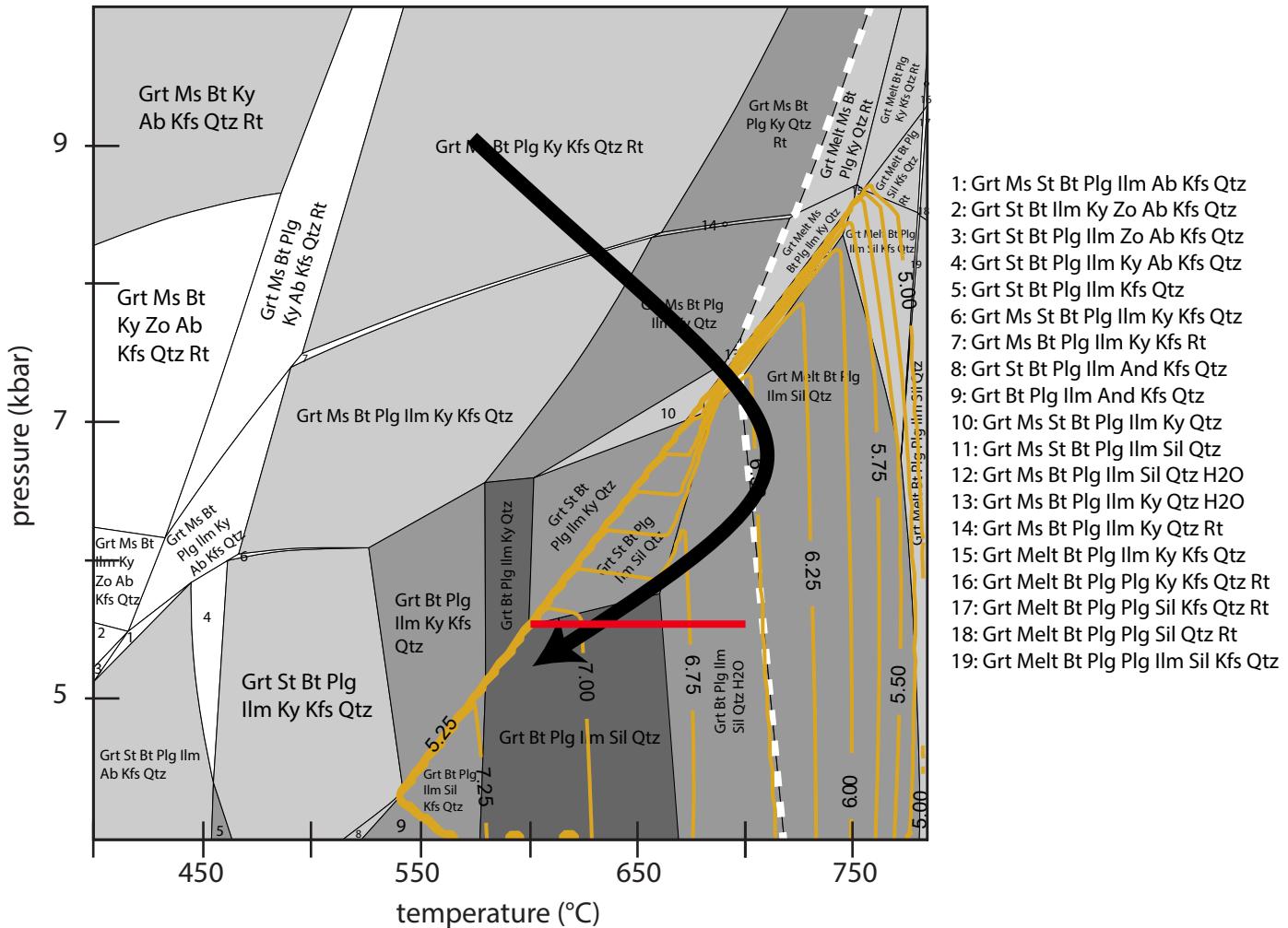
Supplementary Figure 5: P-T phase diagram and constraints on the conditions of plagioclase, muscovite and biotite equilibration. A) Labeled assemblage fields calculated from the composition of sample NT-04 (NT-14-01). The numbered fields refer to the list of assemblages above and to the right. B) P-T conditions of calculated (XMapTools) muscovite Al and K compositions at equilibrium. The positions of these isopleths in P-T space are based on the low water content equilibrium thermodynamic model. C) P-T conditions of the Fe composition (XMapTools) of muscovite at equilibrium. D) Equilibrium conditions of low-Ca plagioclase core compositions. E) Equilibrium conditions for biotite Fe# and Ti compositions. F) Summary of P-T conditions for plagioclase, muscovite and biotite equilibration. All of the isopleths shown here were smoothed using a Gaussian filter.



Supplementary Figure 6: Compositional traverses across garnets in samples A) NT-04 and B) NT-01. We interpret the relatively uniform core compositions to indicate significant diffusional modification of original compositional zoning, making the use of garnet compositions in P-T constraints inappropriate.



Supplementary Figure 7: P-T phase diagram for sample NT-01. Numbered assemblage fields are refer to the list at right. In the lower diagram, the purple contours reflect the modal proportion of garnet at equilibrium. The inset graph shows the modal proportion of garnet along the proposed P-T for sample NT-04 (~10 km along strike). Garnet mode is increasing from 553 to 649°C, so this temperature range is associated with the growth of low Y garnet domains in this sample.



Supplementary Figure 8: P-T phase diagram for sample NT-03 (NT-16-09). This thermodynamic model was calculated with low water contents ($H_2O = 1 \text{ wt\%}$), which likely provides a better match with peak to retrograde conditions. Orange contours represent the modal proportion of sillimanite at equilibrium. The large back arrow shows the proposed P-T path for sample NT-04 (~300m away). Sample NT-03 exhibits large amounts of sillimanite commonly intergrown with biotite. The red bar represents the range of temperature conditions ($650 \pm 50^\circ\text{C}$) where the modal proportion of sillimanite is similar to that seen in thin section and the proposed P-T path would result in sillimanite growth. We associate this temperature range with the age of monazite cores in this sample.